

# **Uncertainty in Hydrologic Modeling and the Curse of Dimensionality**

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# **ACCOUNTING FOR UNCERTAINTY IN HYDROLOGIC MODELLING**

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- **UNCERTAINTY IN MODELLING**

## **Causes or Sources**

- ▶ **natural randomness**
- ▶ **estimation**
- ▶ **scale of representation**
- ▶ **omissions**

- **ACCOUNTING FOR UNCERTAINTY**

## **Approaches**

- ▶ **Monte Carlo simulation**
- ▶ **stochastic equations**
- ▶ **integration**
- ▶ **calculus of propagation**
- ▶ **compensating tricks**



**PASSAGE from the "PUNCTUAL"  
to the "PARCEL"**

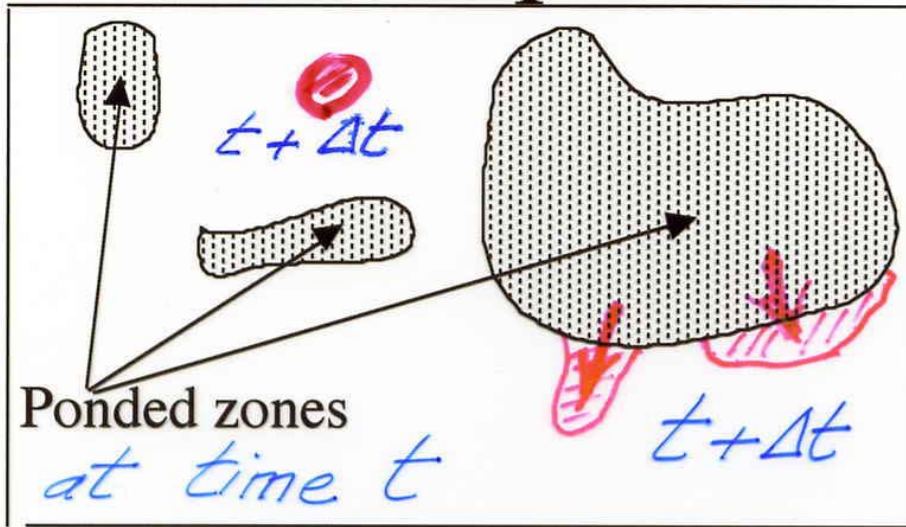
**ENTAILS (ARBITRARY) CHOICE OF**

- **PUNCTUAL, PLAUSIBLE,  
PHYSICAL LAW**
- **PLAUSIBLE LAW OF CHANCE**

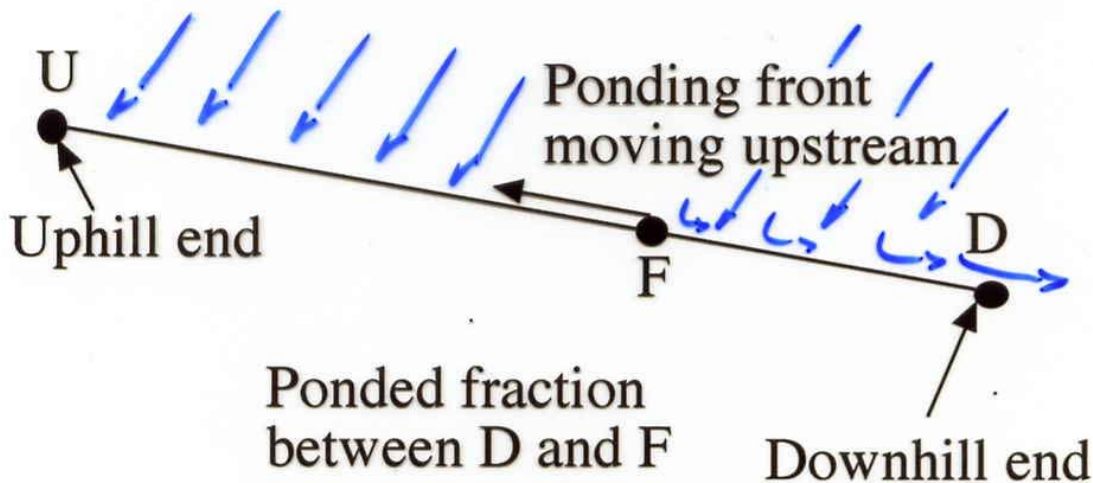
Given "**capacity infiltration rate**" and ponding time formulas at a **point scale**, and laws of **chance** for the distribution of conductivity, one can derive **analytically** the infiltration rate for a plot or a parcel.

In this derivation it is **necessary** to account for runoff from the ponded surfaces that will cause "**run-on**" and thus **opportunity** for infiltration in zones where all the rainfall had infiltrated but some **unfilled capacity** still exists.

# Random pattern



# Organized pattern (such as KDD)



## Typical (possible) formulae

Ponding time : 
$$t_p = \frac{KS_f}{r(r - K)}$$

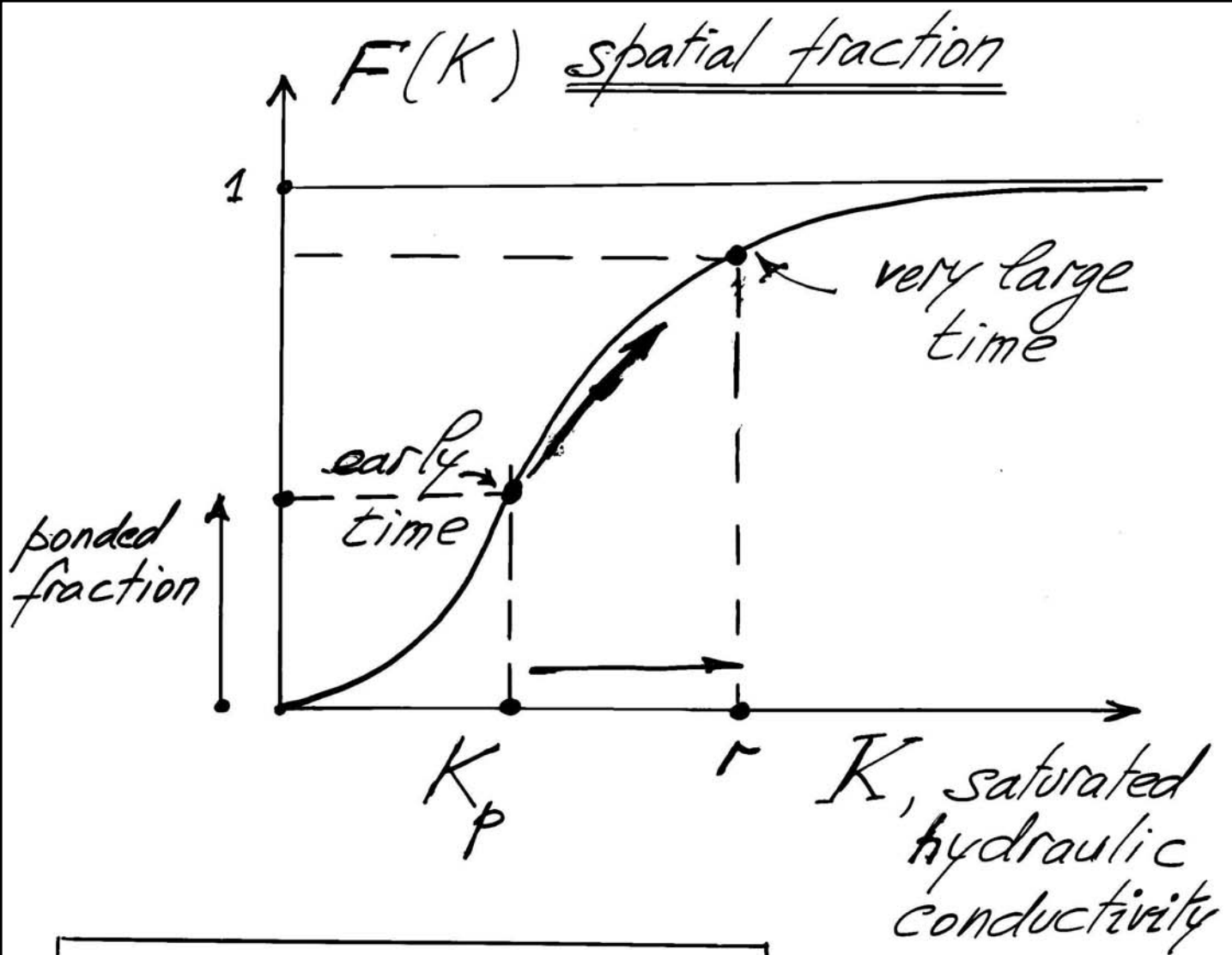
with storage-suction factor:

$$S_f = (\tilde{\theta} - \theta_i) \int_0^{h_{ci}} k_{rw} dh_c = (\tilde{\theta} - \theta_i) \underline{H_{cMi}}$$

and capacity infiltration rate:

$$i_c = K + (r - K)e^{-k(t - t_p)}$$





$$K_p = \frac{r^2 t}{rt + S_f}$$

$S_f$  : storage - suction factor

$(\tilde{\theta} - \theta_i)$   $H_c$   
 natural saturation initial effective capillary drive

## Parcel infiltration rate

Over the ponded fraction

$$\bar{i}_{pfP} = \frac{\int_0^Z i_c(K, t) f(K) dK}{F[Z]}$$

For the entire parcel:

$$i_P = \{1 - F[Z]\}r + \int_0^Z i_c(K, t) f(K) dK$$

The **functional** form  
of infiltration for  
parcel  $\neq$  column



# (SPATIAL) DISTRIBUTION OF CONDUCTIVITY

Density function for  $K$ :  $f(K)$   
cdf:  $F(K)$ .

## "RULES OF THE GAME"

Scenarios as to the fate of the excess rain, i.e. the water that did not infiltrate.

Intantaneous Removal  
(by Microchannels) **IRS**

K Decrease Downhill **KDD**

K Increase Downhill **KID**

Single Huge Uniform  
Column **SHUC**

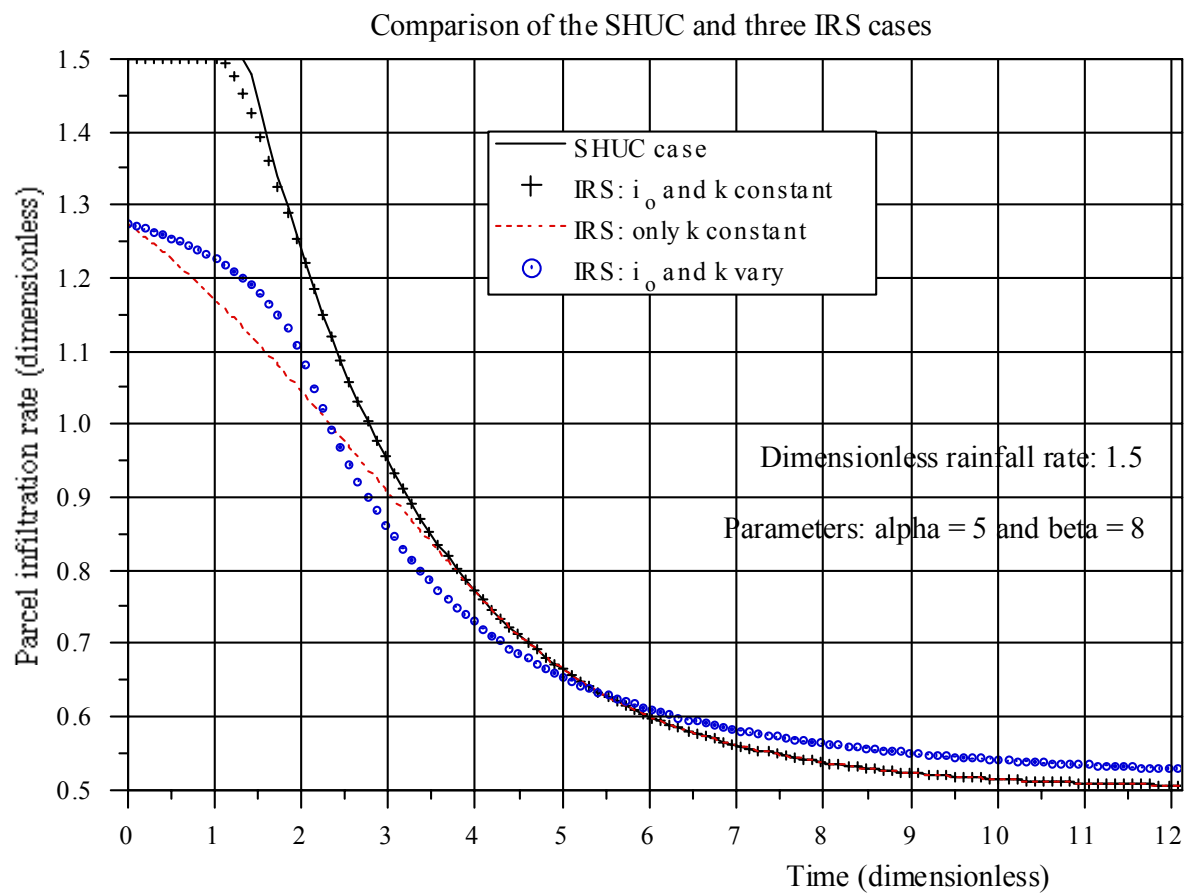


Figure 1. Infiltration rates for the SHUC case and three IRS cases

Comparison between the SHUC and the KID case with  $i_0$  and  $k$  varying

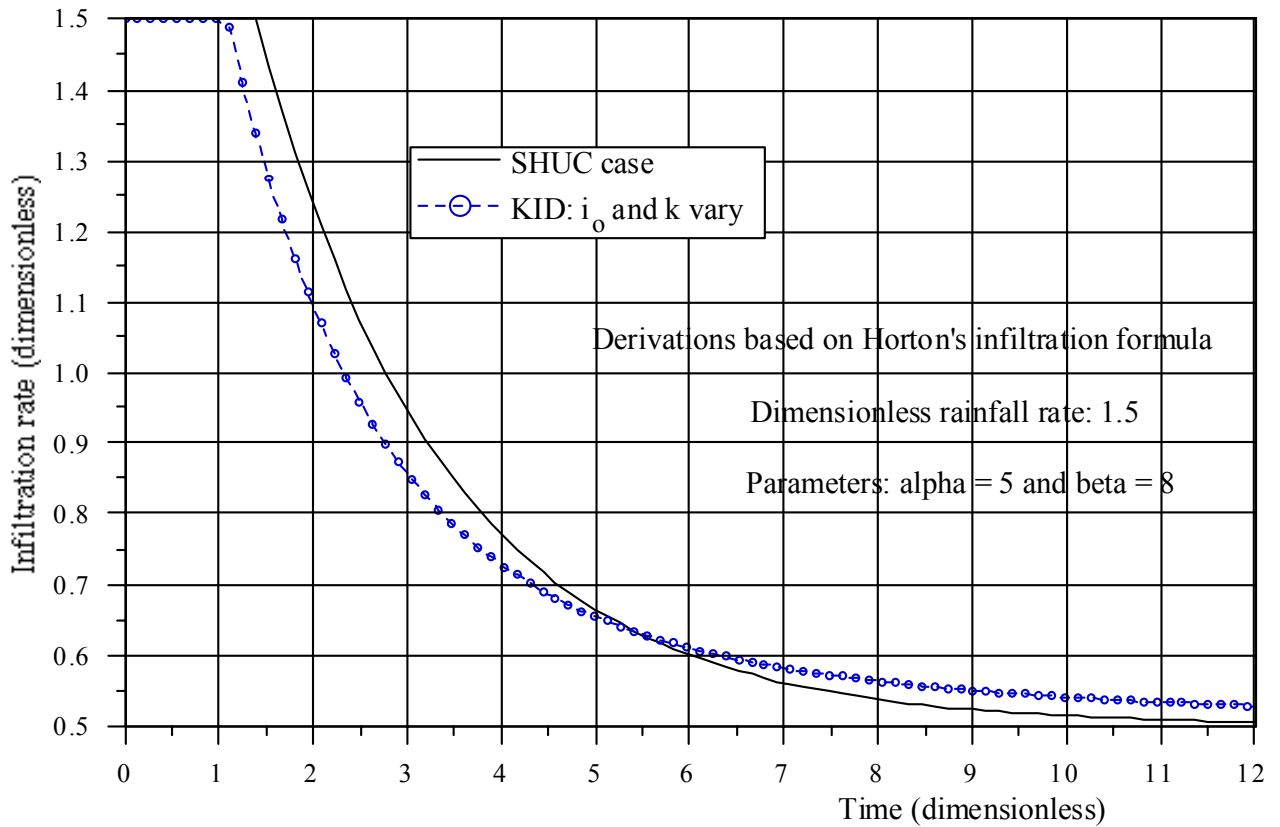


Figure 2. Infiltration rates for the SHUC case and the KID case when all parameters vary.

# CONCLUSIONS

WHICH SCENARIO TO  
USE FOR PARCEL  
INFILTRATION? A MIX.

fraction  $f_\mu$  of the parcel well  
drained by microchannels

$$i_P = f_\mu i_\mu + (1 - f_\mu) i_S$$

More precise derivations will  
**not** significantly improve the  
description of the phenomena as  
opposed to knowing **better** the  
**fraction drained by  
microchannels versus that  
subject to run-on.**



## **Stochastic** (Groundwater) Equation. #1

**Classical deterministic** Boussinesq equation

$$S \frac{\partial H_r}{\partial t} - \text{div.}[T_r(\text{grad.} H_r)] = q$$

with usual symbols but with the **added** subscript **r** to indicate that such a subscripted variable is actually a **random** variable. To simplify let us only consider the **steady-state** case in just **one** dimension and with **no** recharge or withdrawal rate within the domain.

$$\frac{\partial}{\partial x} [T_r(\frac{\partial H_r}{\partial x})] = 0$$

## Stochastic (Groundwater) Equation. #2

$$\frac{\partial}{\partial x} \left[ T_r \left( \frac{\partial H_r}{\partial x} \right) \right] = 0$$

Define:  $H_r(x) = H(x) + h(x)$

$H(\mathbf{x})$  is the **mean** value of  $H_r(\mathbf{x})$ .

The expected value of  $h$  is zero.

Similarly:  $T_r(x) = T_L(x)e^{\theta(x)}$

The expected value of  $\theta$  is zero.

The **random** variable  $\theta$  is defined

by its variance:  $\sigma_{\theta}^2(\mathbf{x})$  and

covariance:  $C_{\theta, \theta'}(x, x')$ .

**Given** values and functions.

## Stochastic (Groundwater) Equation. #3

$$\frac{\partial}{\partial x} \left[ T_r \left( \frac{\partial H_r}{\partial x} \right) \right] = 0$$

Substitution (with previous definitions) :

$$\frac{\partial}{\partial x} \left[ T_L e^{\theta} \frac{\partial H}{\partial x} + T_L \left( \frac{\partial h e^{\theta'}}{\partial x} \right)_{x'=x} \right] = 0$$

Taking **expectation**, defining:  $T_A = T_L e^{\frac{1}{2} \sigma_{\theta}^2}$

$$\frac{\partial}{\partial x} \left[ T_A \frac{\partial H}{\partial x} + T_A \left( \frac{\partial C_{h\theta'}}{\partial x} \right)_{x'=x} \right] = 0$$

Problem? One equation and **two** unknowns:

$$H(\mathbf{x}) \text{ and } C_{h\theta'}(\mathbf{x}, \mathbf{x}')$$

## Stochastic (Groundwater) Equation. #4

Premultiply the stochastic equation

$$\frac{\partial}{\partial x} \left[ T_L e^{\theta} \frac{\partial H}{\partial x} + T_L \left( \frac{\partial h e^{\theta'}}{\partial x} \right)_{x'=x} \right] = 0$$

by **random** variable  $\theta' (x')$  and take **expectation** leading eventually to:

$$\frac{\partial}{\partial x} \left[ T_A C_{\theta, \theta'} \left\{ \frac{\partial H}{\partial x} + \left( \frac{\partial C_{h\theta'}}{\partial x} \right)_{x'=x} \right\} + T_A \left( \frac{\partial C_{h\theta'}}{\partial x} \right) \right] = 0$$

Now this is a system of **two** equations for **two** unknowns for all  $x$  and  $x'$  within the domain of interest.

Typical choice for:  $C_{\theta, \theta'} = \sigma_{\theta}^2 e^{-|x-x'|/D}$